

Is the ozone layer recovering ?

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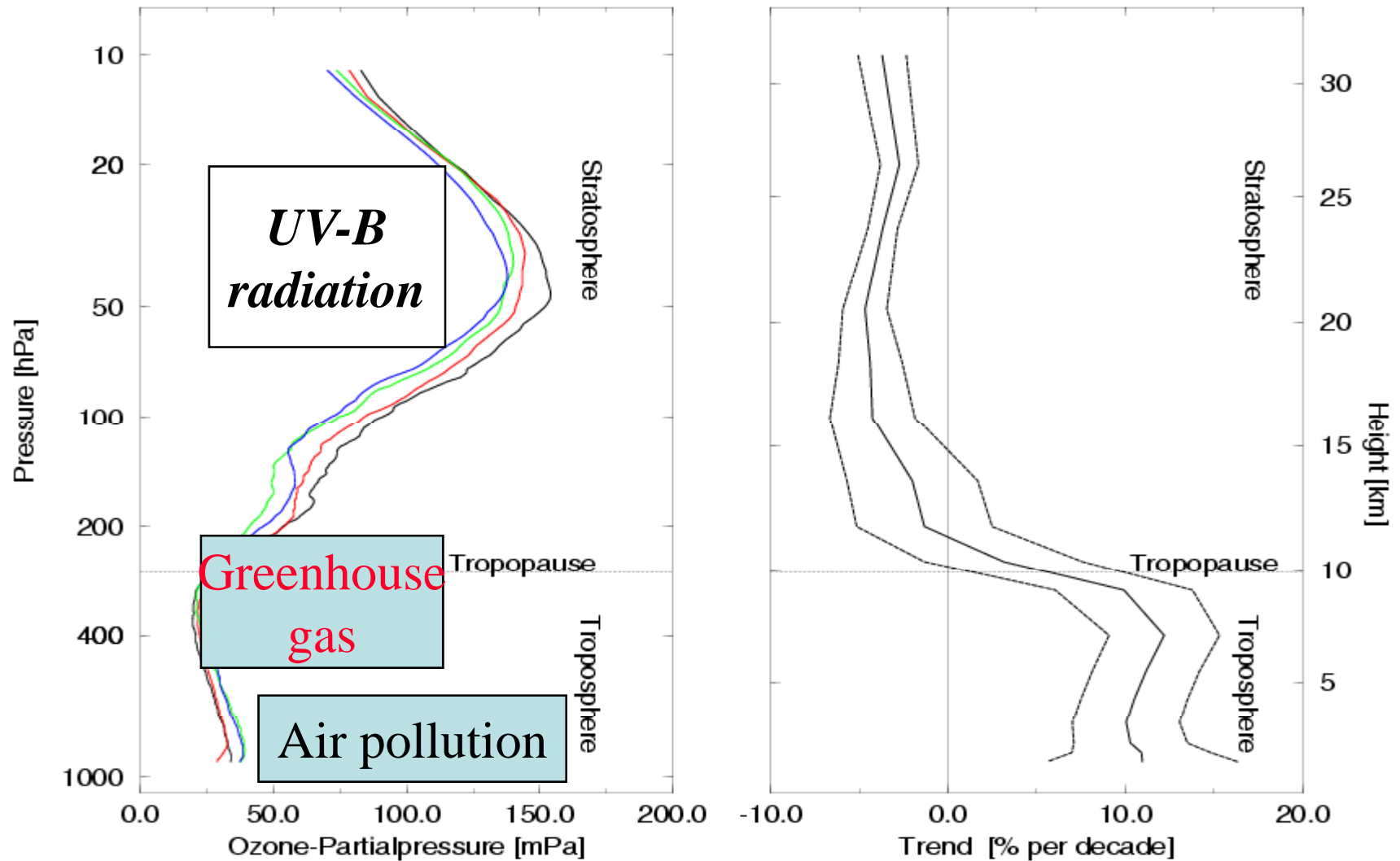
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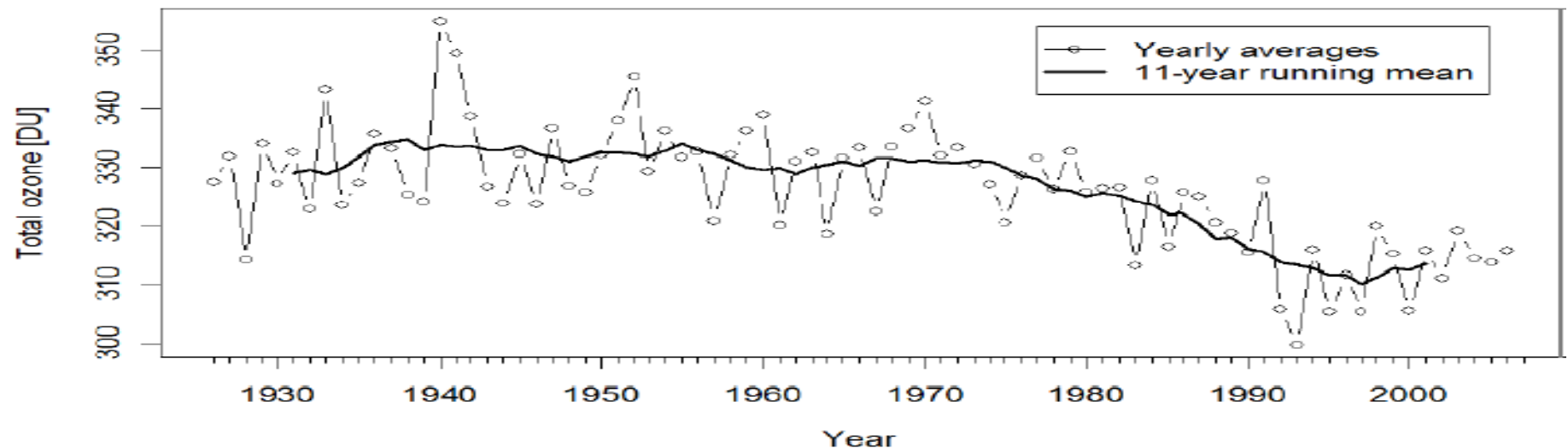
1. Introduction

Measurements of ozone sondes of Payerne (Switzerland). Black: 1970; red: 1980; green: 1990; blue: 2000



Swiss long-term ozone measurements (MeteoSwiss since 1988)

1. Longest total ozone series of the world (Dobson spectrophotometry), homogenised
2. First Umkehr measurements (1930), continuous measurements since 1956
3. Ozone sonde measurements since 1969 (Payerne, Swiss Plateau)



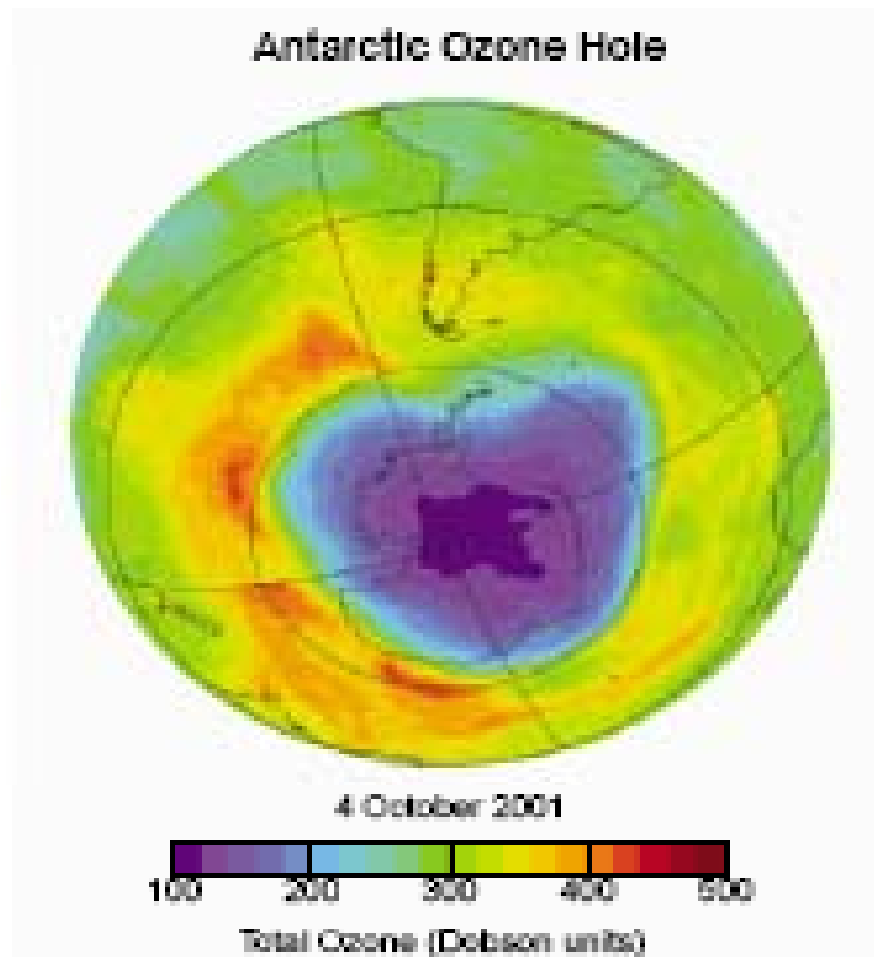
Short history: 1970s

Anthropogenic stratospheric ozone depletion (since early 1970s):

- H. Johnston (1971), P. Crutzen: Ozone depletion by Super Sonic Transport (NO_x) (?)
- R. Stolarski, R.J. Cicerone (1974): Ozone depletion by Chlorine radicals
- M.J. Molina, S. Rowland (1974): Ozone depletion by CFCs

(Ozone Depleting Substances (ODS): CFCs, halones, HCFCs)

Surprise: Farman et al., 1985: Discovery of Antarctic Ozone hole



Second part of 1980s

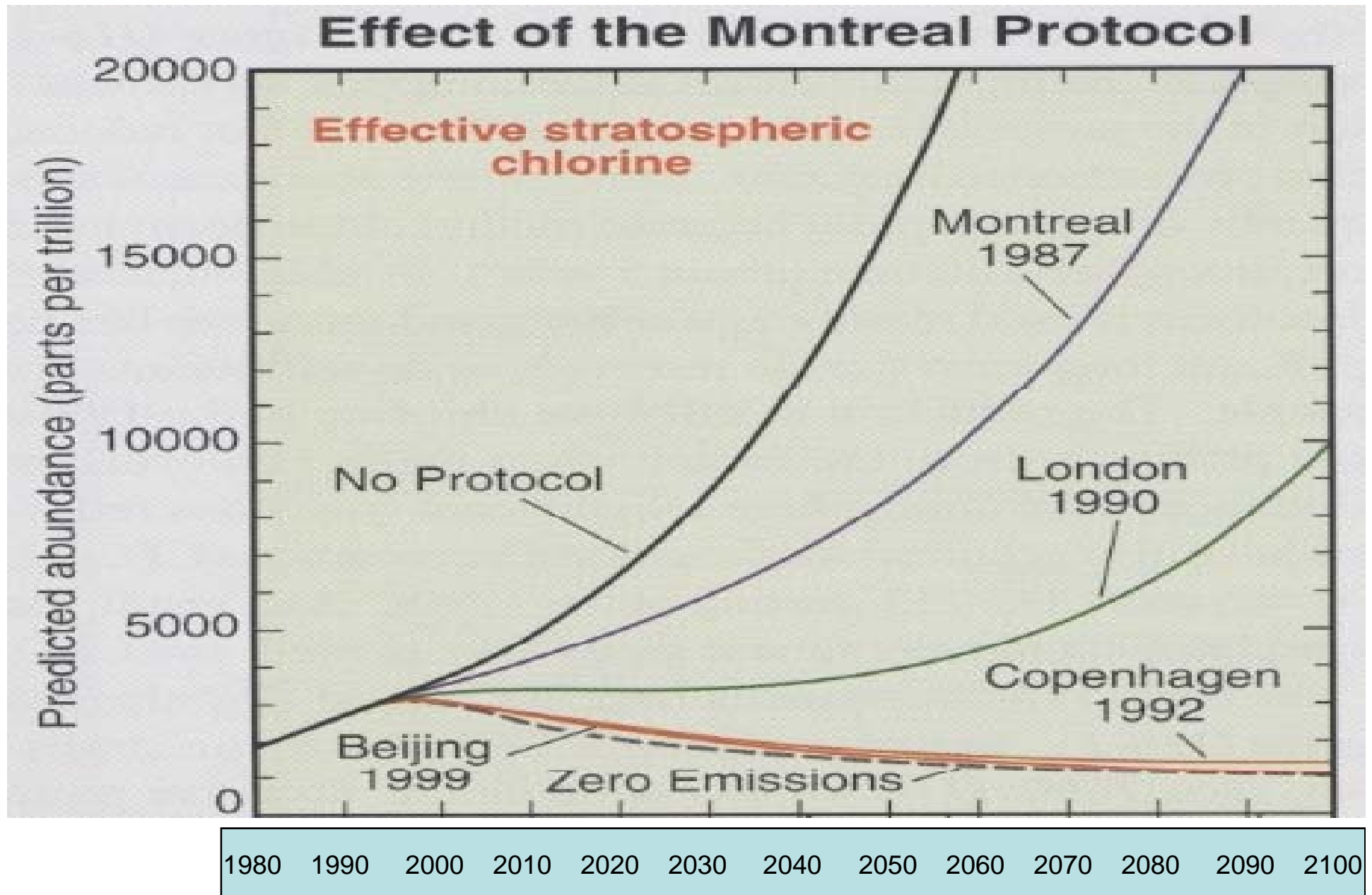
- 1988: Explanation of ozone hole by (anthropogenic) CFCs (halones) (heterogeneous chemistry)
- 1988: Publication of International Ozone Trend Panel Report: Significant decrease in (winter) ozone at Northern midlatitudes (multiple regression analysis)

2. Global Regulation

- 1985: Vienna Convention
- 1987: Montreal Protocol
- Several amendments and adjustments
- Quantity for (chemical) ozone depletion of ODS: **EESC** (Equivalent Effective Stratospheric Chlorine): Weighting over release and ozone depletion of individual ODS

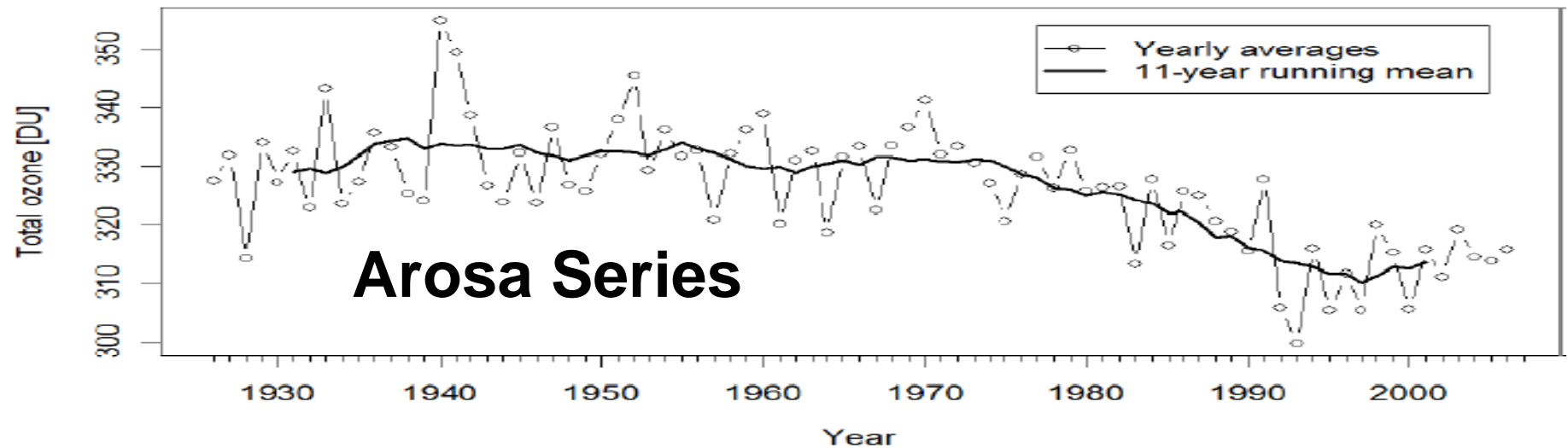
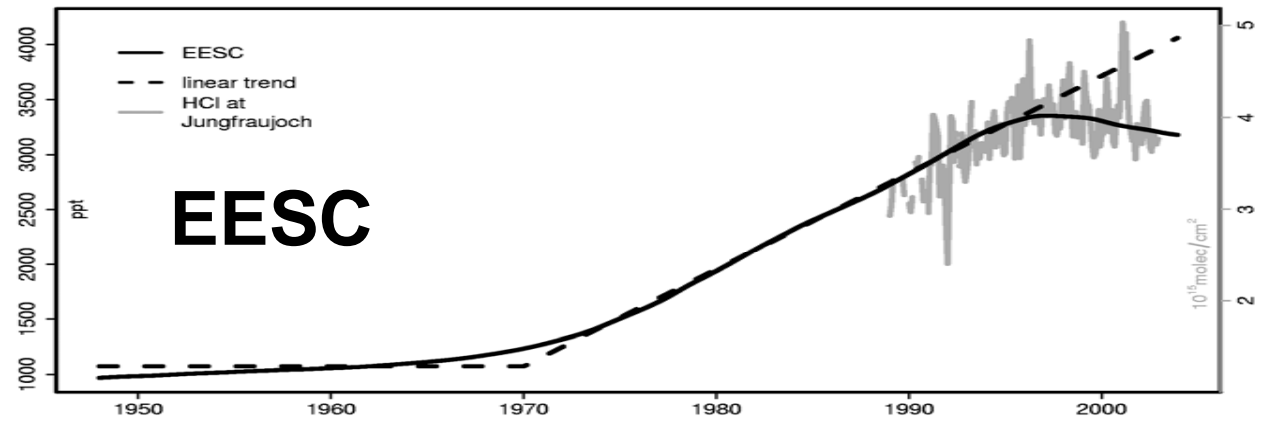
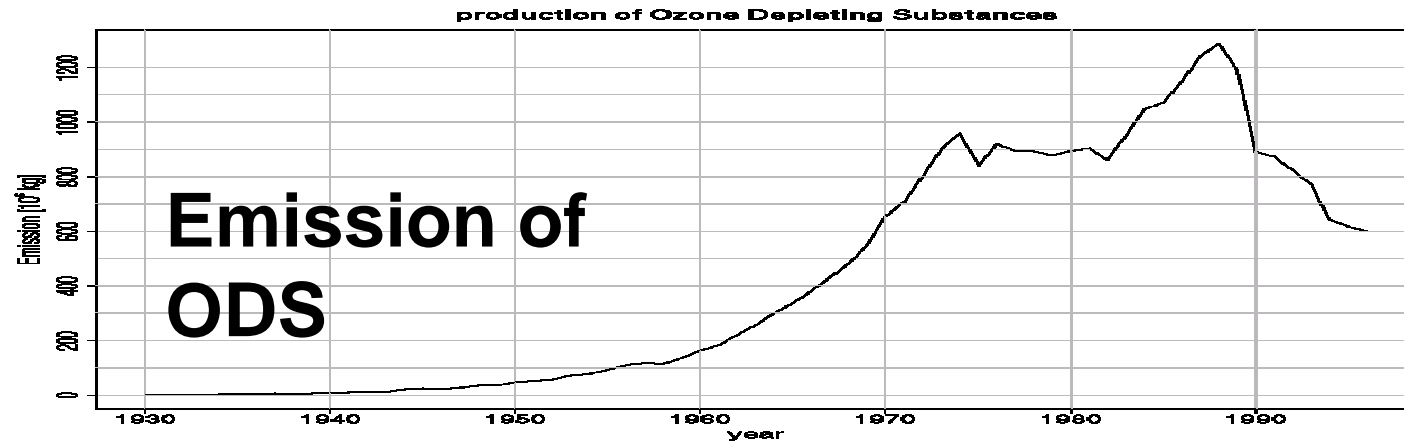
EESC for mid-latitudes

http://www.wmo.ch/web/arep/reports/ozone_2002/q&as.pdf, Seite Q.29



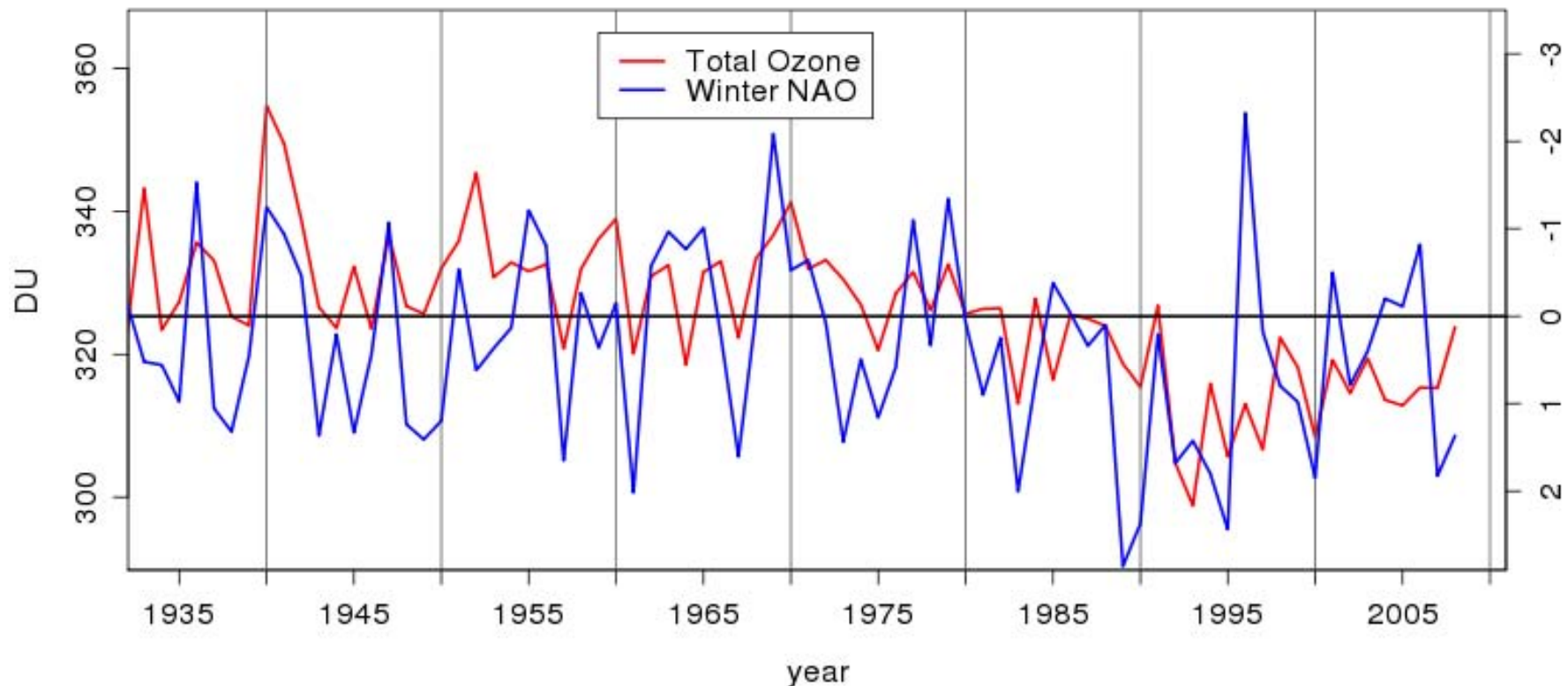
3. Problem of documentation of ozone shield recovery

- Illustration by the total ozone series of Arosa: started in 1926 and continued since 1988 by MeteoSwiss:
- Problem: Attribution of ozone evolution to individual processes (in addition to EESC)



- Further processes affecting mid-latitude ozone
- Violent Volcanic eruptions (Pinatubo, June 1991)
 - (Long-term) climate Variability: Strong correlation between (NAO(AO index and winter total ozone values at Arosa (Appenzeller et al., 2000, updated by J. Mäder), Brewer Dobson Circ., polar ozone depletion

Total ozone at Arosa (Switzerland) versus Winter NAO (Hurrell)



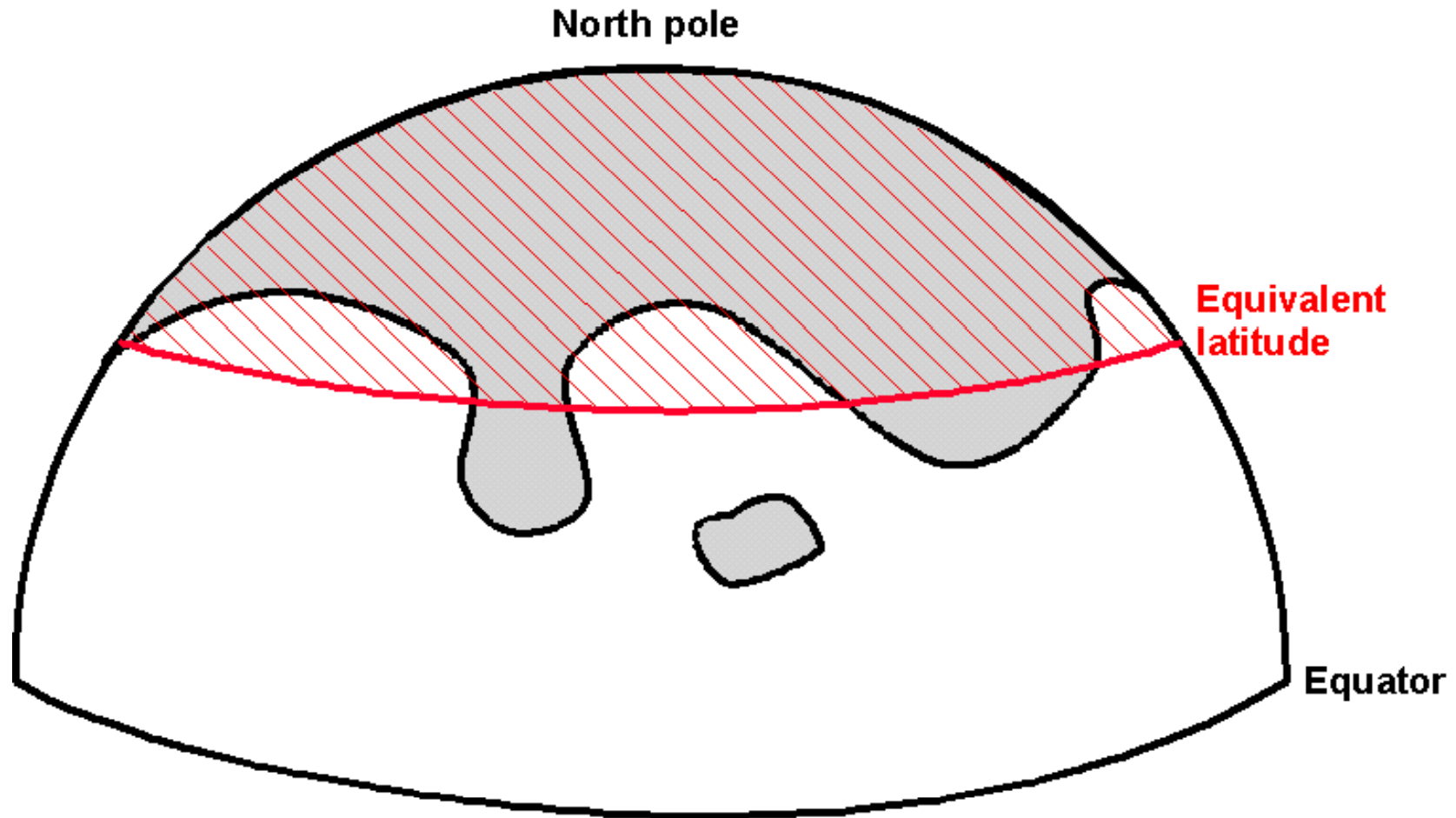
4. CATO ozone data sets

CATO (Candidoz Assimilated Three-dimensional Ozone)
(EU-project: CANDIDOZ: Chemical and Dynamical
Influences on Decadal Ozone Change): Principle:

- Measurements: Satellite total ozone data (since 1979) of NIWA data set (G. Bodeker): Composite of different satellite total ozone measurements normalized by ground-based Dobson measurements
- Assimilation technique: Kalman filtering
- ECMWF (ERA-40) for PV at 16 pot. temp. levels
- Tropospheric ozone residual substracted
- Assimilation technique (PV) not suitable for upper stratospheric ozone: In one version satellite (SBUV) measurements used

(see Brunner et al., J. Geophys Res., 2006)

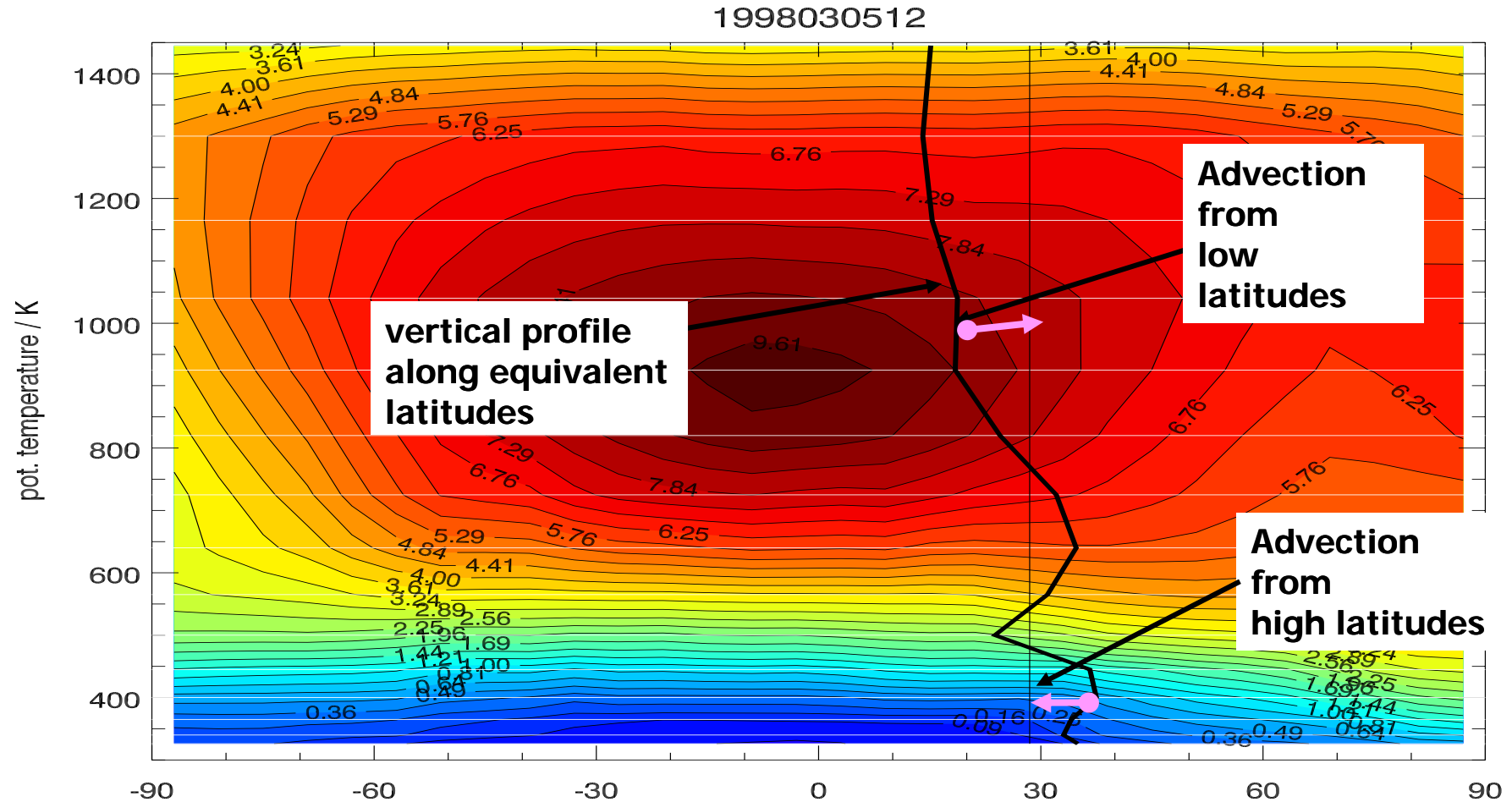
CATO based on equivalent latitude/theta coordinates



$$A = 2 p (1 - \sin f_E) \text{ in spherical radian}$$

Reconstruction method

Vertical column in equivalent latitude – θ space



P = tropospheric residual < 320 K

$$\Omega(\lambda, \varphi) = P(\lambda, \varphi) + a \cdot \int_{\theta_0}^{\infty} d\theta (-\partial p / \partial \theta) \cdot \chi(\varphi_{eq}(\lambda, \varphi, \theta), \theta)$$

5. Multiple regression analysis of CATO

Multiple linear regression model

$$Y_t = a + b(t) + \sum^N c_j X_j(t) + e(t)$$

t: number of month since start of record (t=1: January 1979)

Y_t : Monthly mean total ozone (or ozone partial pressure)

a: seasonally varying intercept (offset) of ozone time series

b: *Trend term: Two hockey stick method*

(or EESC: Equivalent effective stratospheric chlorine)

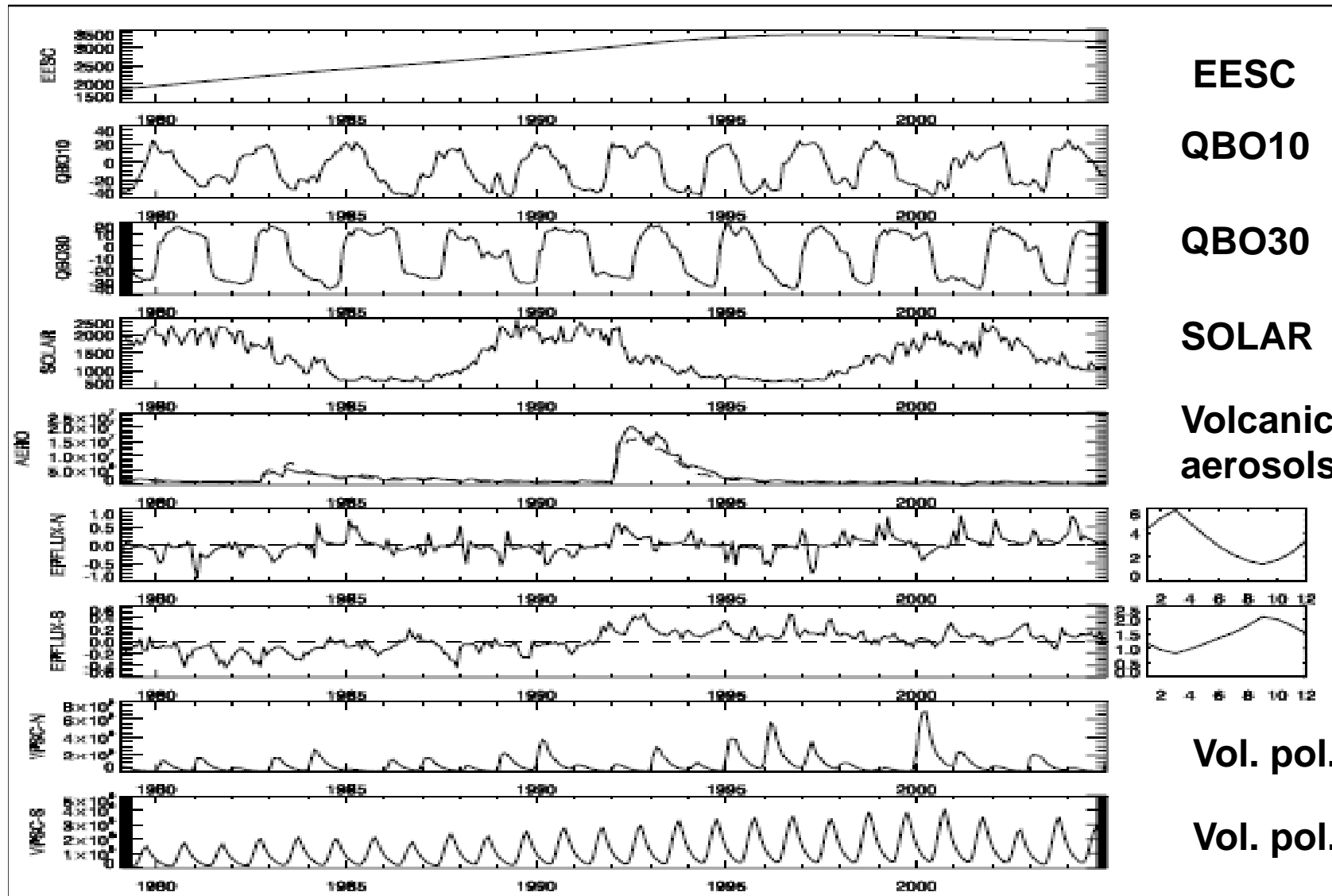
$c_j X_j(t)$: time series of expl. Var. ($j= 1, N$), (seasonally) var. coef

$e(t)$: residual variations (not described by model)

a, b, c: depend on month of year, described by 12-month and 6-month harmonic series:

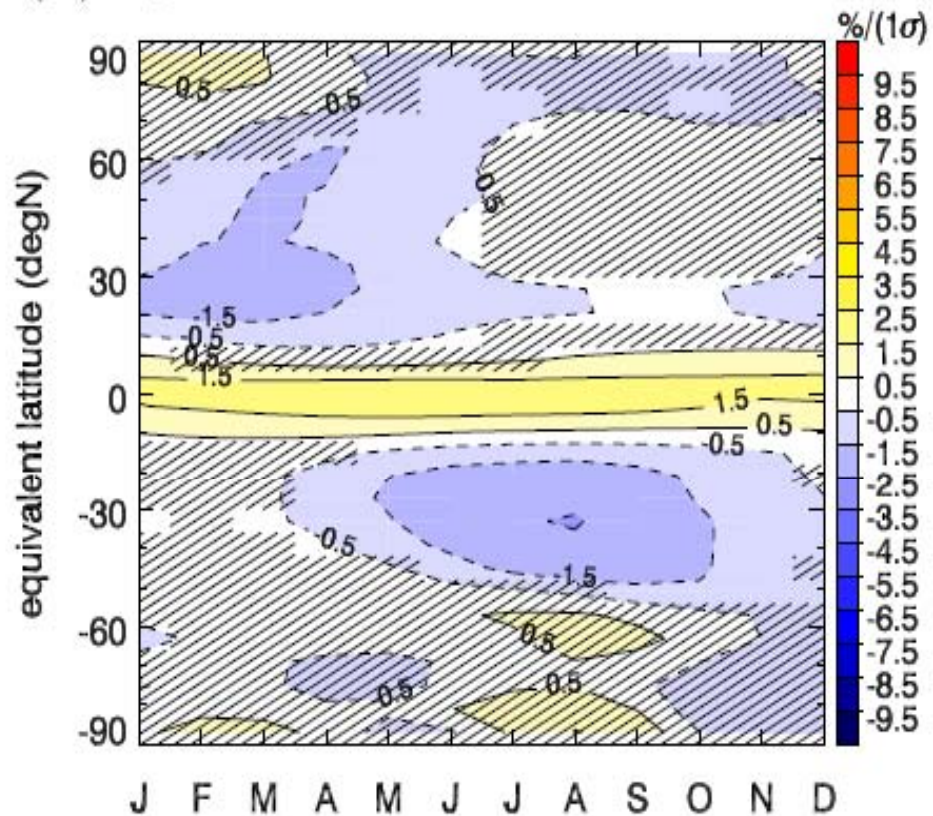
$$c_j(t) = c_j(1) + \sum^2 (c_{j,2k} \cos(2pkt/12) + c_{j,2k+1} \sin(2pkt/12))$$

Used explanatory variables:

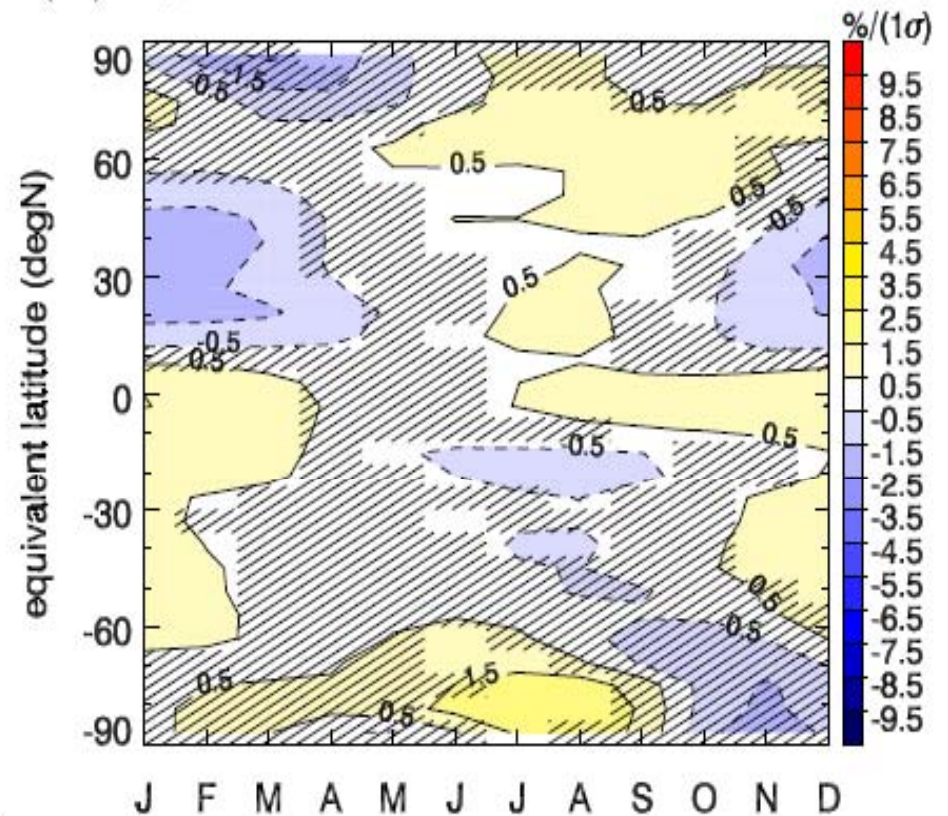


Contribution of QBO(30hPa) and QBO(10hPa) to total ozone variability (funct. of season and equiv. latitude): Regression coefficients multiplied by 1σ of each proxy time series and then divided by 1979-2004 mean ozone distribution (% change in tot. O_3 for 1σ increase in proxy)

(a) QBO at 30 hPa

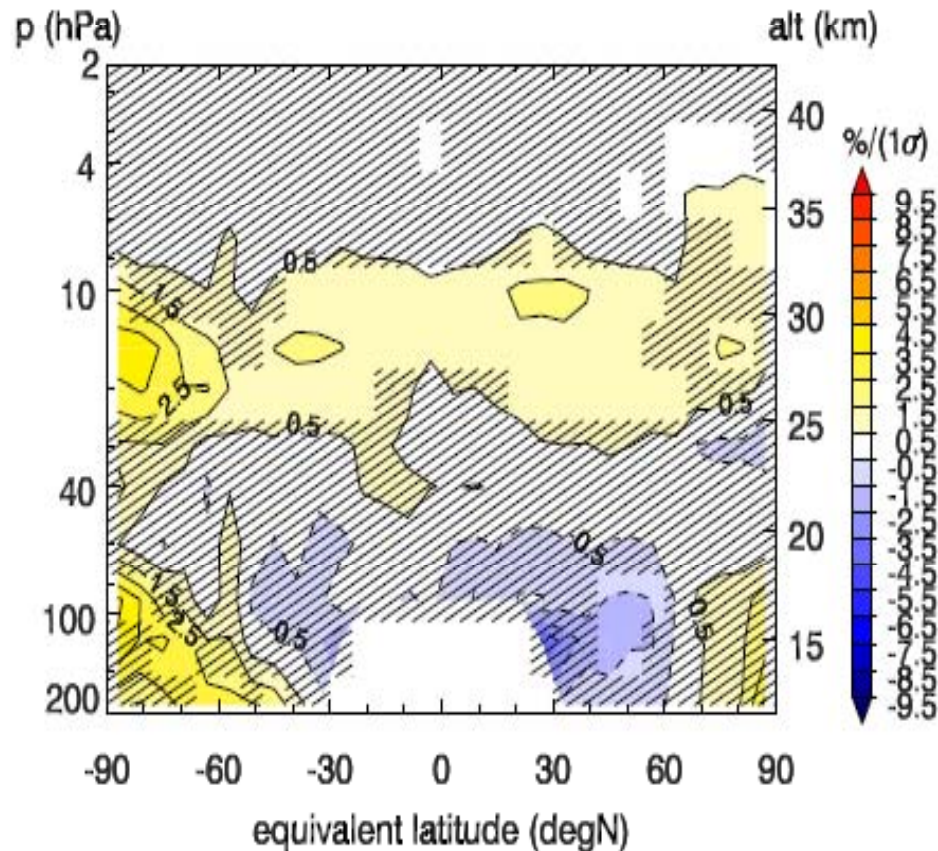


(b) QBO at 10 hPa

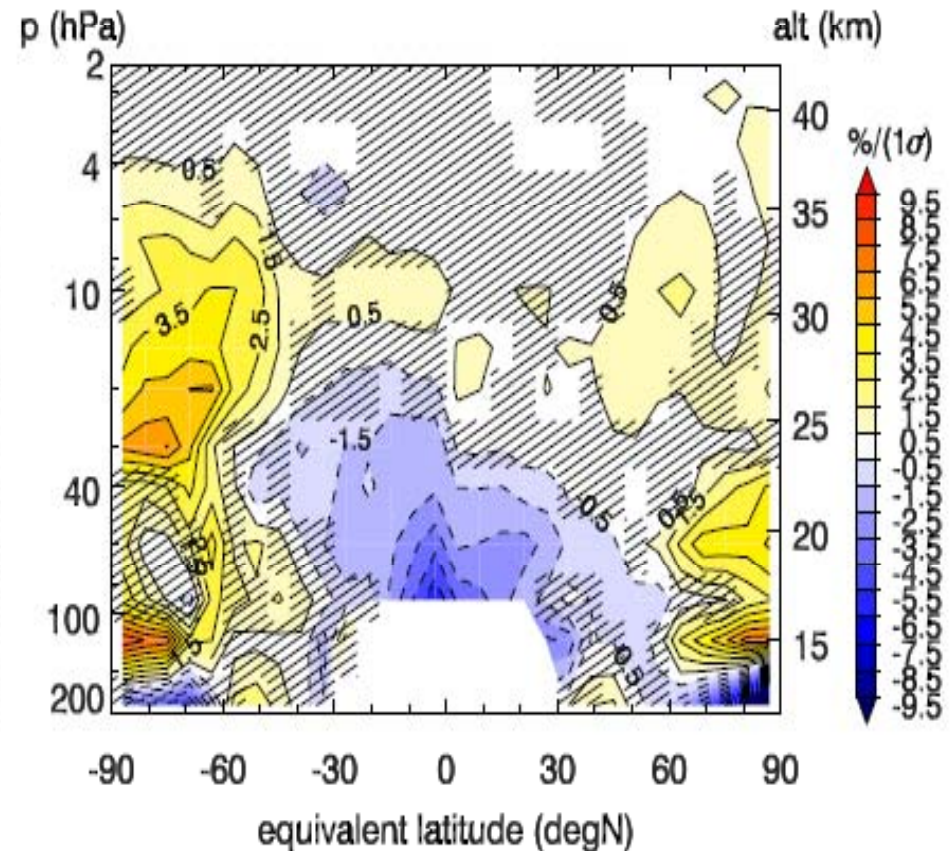


Annual average contribution to variability in ozone partial pressure (altitude/equivalent latitude)
(% change in ozone concentration for 1σ increase in proxy)

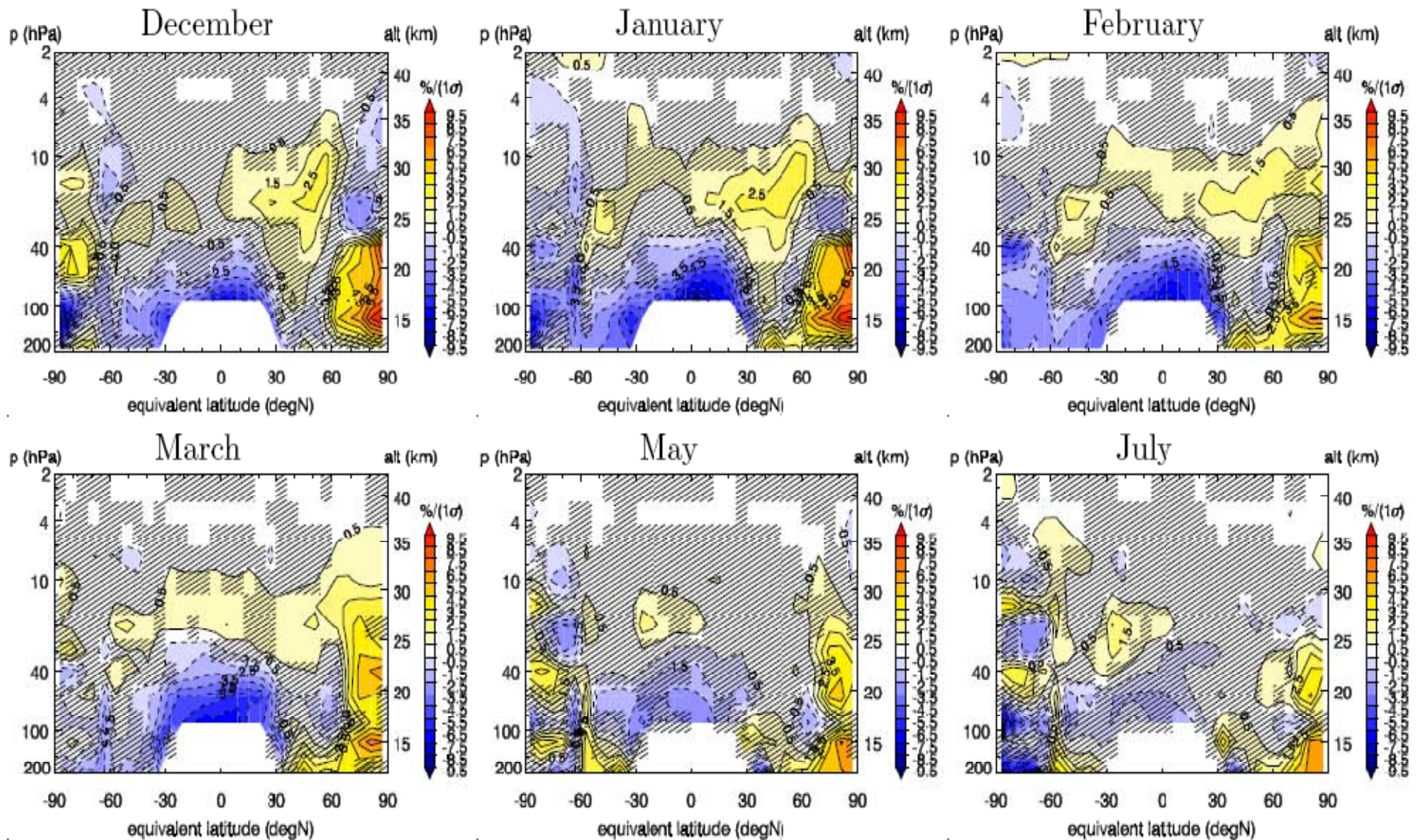
(d) Solar cycle



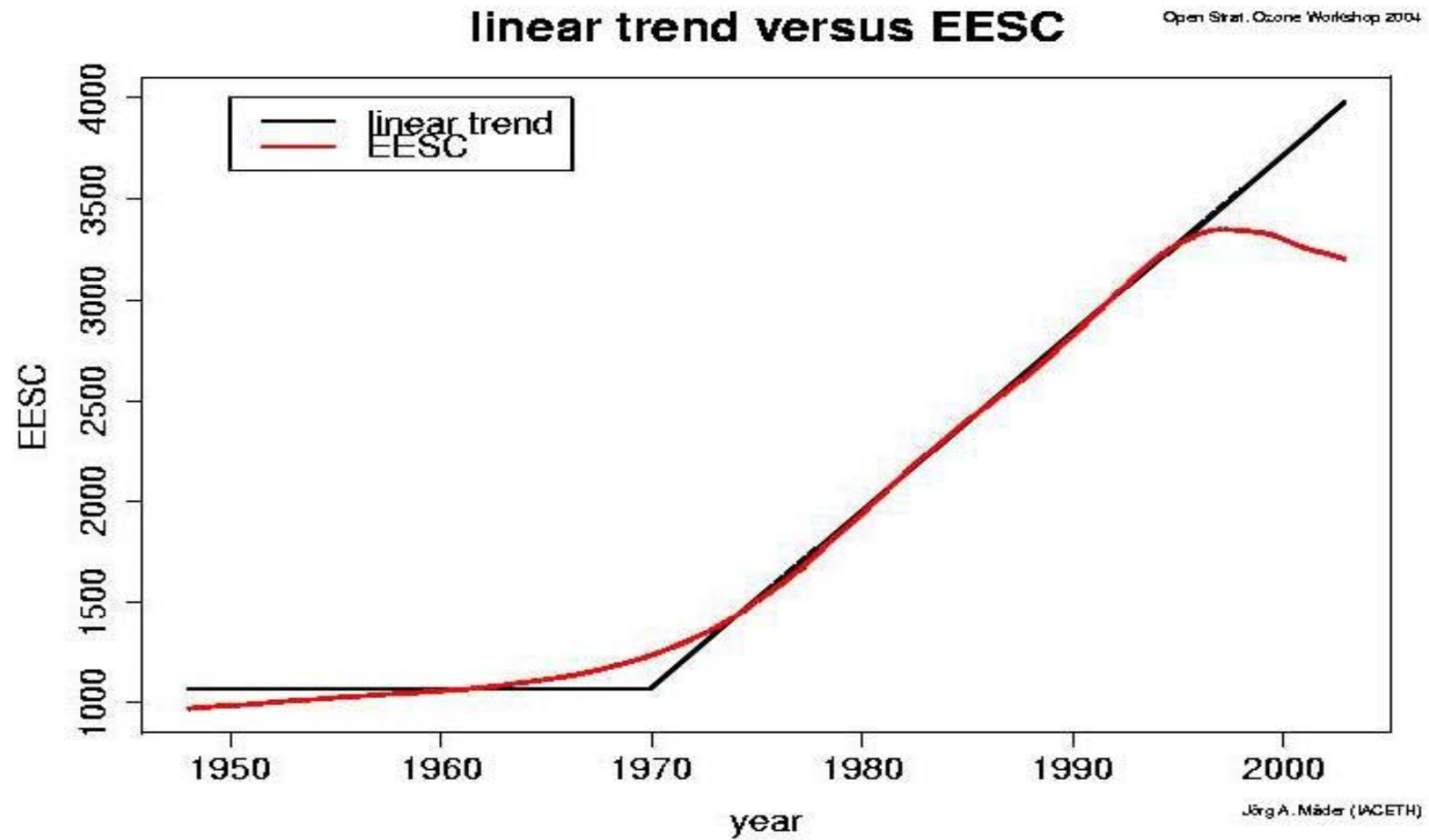
(e) EP flux



Sequence of hemispheric EP-flux on global stratospheric ozone distribution



“Turn-around”: Two hockey stick (Reinsel et al., 2005)



Mathematical description (Reinsel et al.)

$$Y_t = \mu + S_t + \omega_1 X_{1t} + \omega_2 X_{2t} + \dots + \gamma_i Z_{i,t} + N_t$$

t : month (1, .., T), period: 1978-2004

Y_t : Monthly mean total ozone; μ : baseline constant

S_t : Seasonal component (linear fit of sin/cos functions)

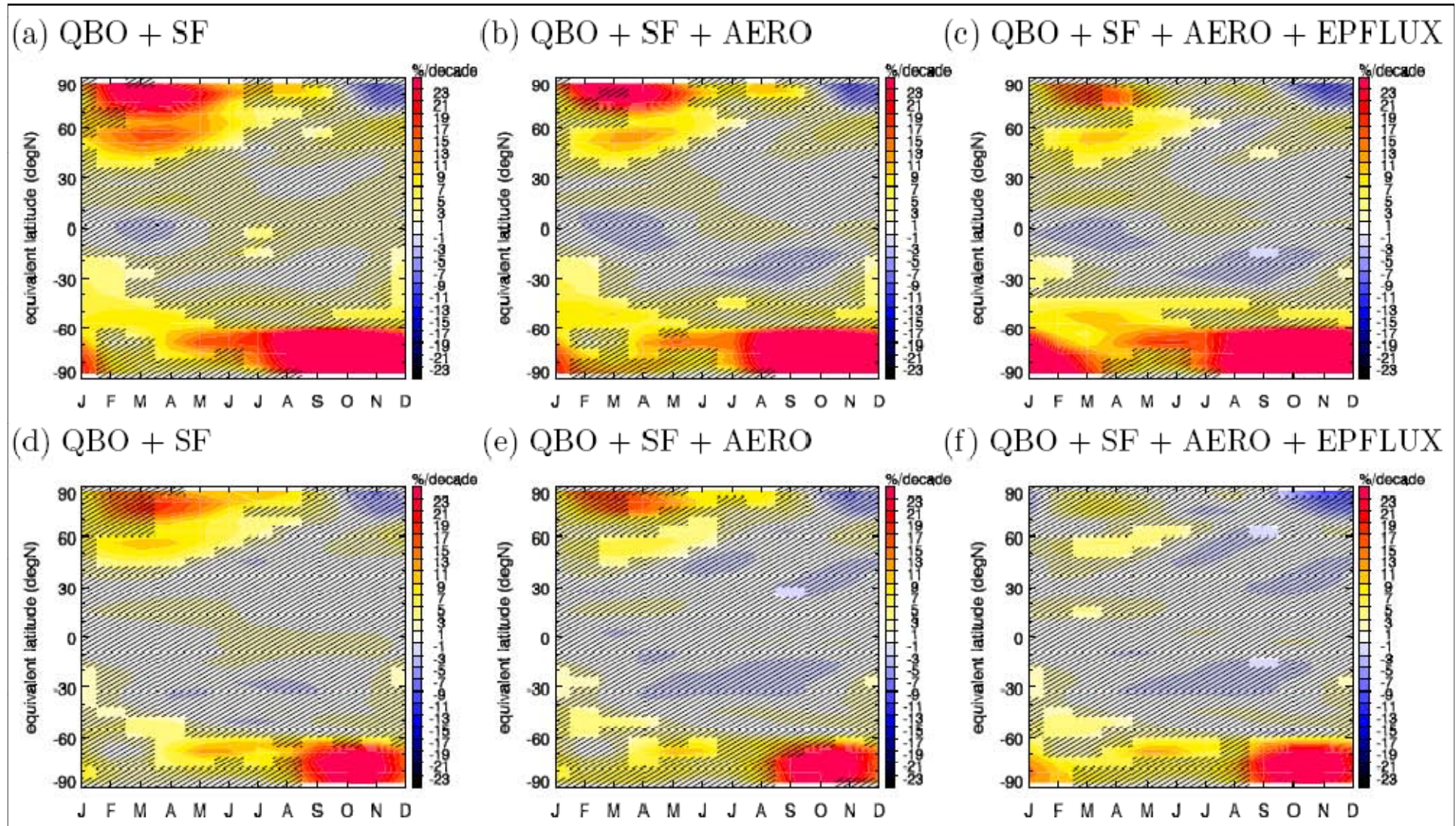
ω_1 : **linear (decreasing) trend**, beginning at $t=0$:
effect of ODS

ω_2 : **change: linear additional upward trend**,
starting from 1996: ***effect of Montreal Protocol***

$N_t = \rho N_{t-1} + \varepsilon$: autoregressive noise term;

ε : independent random variables

Key results (1979-2004) (a)-(c): Changes in ω ; (d)-(f): upward trend



Conclusions from Brunner et al., ACP, 2006

- Equivalent latitude coordinates partially compensate for short term variability
- Polar ozone depletion can not be separated from changes in Brewer Dobson circulation
- Results regarding “recovery” depends on used explanatory variables
- Using data until 2004: Only marginal sign of effect of Montreal Protocol regulation

Antarctic ozone hole

<http://www.wmo.int/web/arep/gawozobull06.html>

No indication of recovery: Extent of Antarctic ozone depletion depends on meteorological cond.

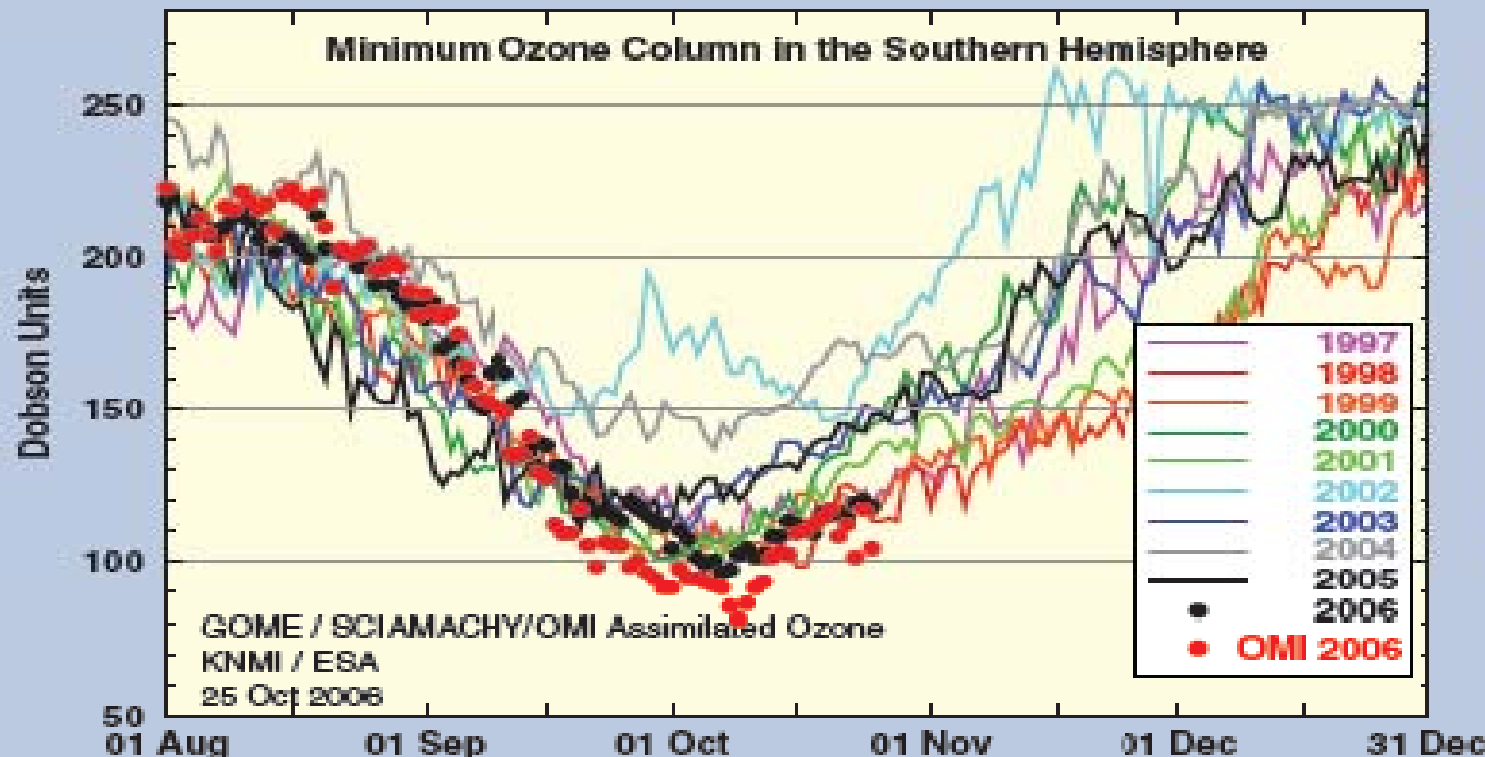


Figure 4. Daily minimum total ozone columns in the Southern Hemisphere as observed by GOME, SCIAMACHY and OMI from 1997 to now. The black dots show the SCIAMACHY observations for 2006. The red dots show the OMI observations for 2006. The SCIAMACHY data now show minimum ozone columns down to 110 DU. The OMI data show somewhat lower minimum ozone columns than SCIAMACHY since OMI looks deeper into the vortex and now sees minimum ozone columns around 100 DU. The plot is provided by the Netherlands Meteorological Institute (KNMI).

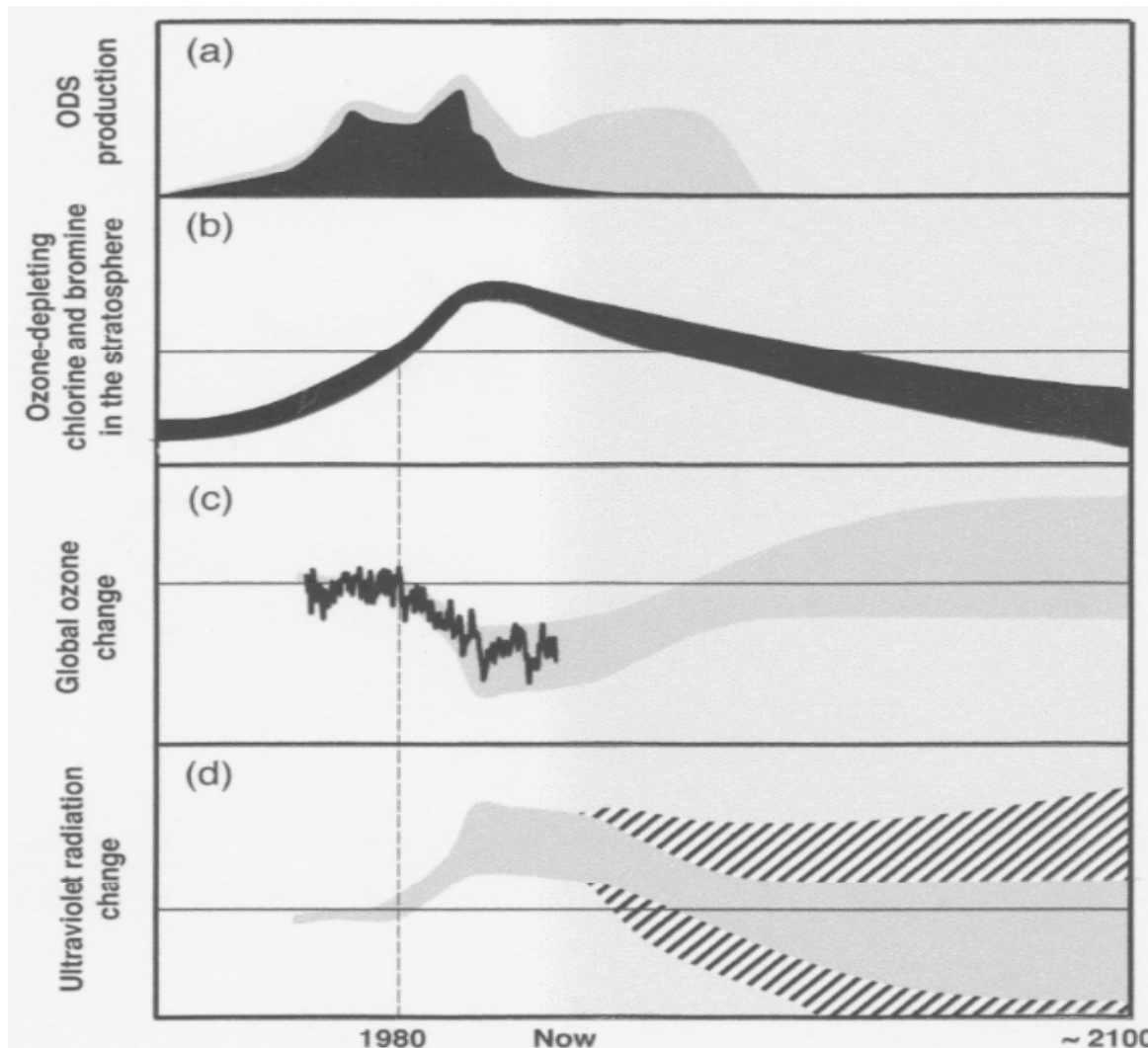
6. Conclusions

- Record low values in 1992, increase since 1993: mainly attributable to Pinatubo eruption (1991) and (long-term) climate variability, etc.
- No signs of recovery of Antarctic ozone (in agreement with expectation)
- Attribution of midlatitudinal changes to processes (Montreal Protocol (1987) vs. others) still challenging
- Continuation of high quality ground-based and satellite ozone measurements important

7. Expectation

(Executive Summary, Scientific Assessm. of Ozone Depletion:
2006, WMO/UNEP, 18. 8. 2006

(http://www.wmo.int/web/arep/reports/ozone_2006/exec_sum_18aug.pdf)



(a) Produktion of ODS
(black: CFCs and
halones; grey:
HCFCs (used as
replacement)

(b) Effect of ODS on
midlatitudinal
stratospheric ozone
(EESC)

(c) Ozone changes 60°S-
60°N (black:
Measurements; grey:
numerical modles)

(d) Effect of ozone
depletion on
erythemal UV (grey
according (c),
hattched: including
additional processes)

(Possible) future development

- Montreal Protocol effective to reduce (anthropogenic) ODS emissions
- Recovery of global ozone layer expected
- Future challenges: Effect of climate change on stratospheric ozone
- “Super recovery” caused by stratospheric temperature decrease ?
- Intensification of Brewer-Dobson circulation ?
- Increased transport of stratospheric ozone to troposphere ? Effect on tropospheric ozone budget ?